

Time-dependent, axisymmetric model phrased in R space

- Hydrostatic and gradient balance above PBL
- Moist adiabatic lapse rates on M surfaces above PBL
- Boundary layer quasi-equilibrium
- Deformation-based radial diffusion

Potential Radius:

$$\frac{f}{2} R^2 \equiv M = rV + \frac{f}{2} r^2 \quad (1)$$

Local energy conservation:

$$\frac{1}{r_b^2} - \frac{1}{r_t^2} = \frac{-2(T_s - T_t)}{f^2 R^3} \frac{\partial s^*}{\partial R} \quad (2)$$

Differentiate in time:

$$\frac{1}{r_b^3} \frac{\partial r_b}{\partial \tau} - \frac{1}{r_t^3} \frac{\partial r_t}{\partial \tau} = \frac{(T_s - T_t)}{f^2 R^3} \frac{\partial}{\partial R} \frac{\partial s^*}{\partial \tau} \quad (3)$$

Mass continuity:

$$ru = \frac{\partial \psi}{\partial p},$$

$$r\omega = -\frac{\partial \psi}{\partial r}$$

Transform to potential radius coordinates:

$$ru = r \frac{dr}{dt} = r \left[\frac{\partial r}{\partial \tau} + \frac{\partial r}{\partial R} \frac{dR}{dt} + \frac{\partial r}{\partial P} \frac{dP}{dt} \right] = \frac{\partial \psi}{\partial p},$$

$$\rightarrow r \frac{\partial r}{\partial \tau} = \frac{\partial \psi}{\partial p} - r\omega \frac{\partial r}{\partial P} = \frac{\partial \psi}{\partial p} + \frac{\partial \psi}{\partial r} \frac{\partial r}{\partial P} = \frac{\partial \psi}{\partial P}$$

$$\boxed{\frac{\partial r^2}{\partial \tau} = 2 \frac{\partial \psi}{\partial P}}$$

Define ψ_0 as streamfunction at top of boundary layer and use simple finite difference in vertical:

$$\begin{aligned} \frac{\partial r_b^2}{\partial \tau} &\cong 2(\psi_0 - \psi), \\ \frac{\partial r_t^2}{\partial \tau} &\cong 2\psi \end{aligned} \tag{4}$$

PBL flow:

$$\cancel{r \frac{\partial \cancel{r}}{\partial \tau}} + r \frac{\partial r}{\partial R} \frac{dR}{dt} = \frac{\partial \psi}{\partial P}$$

Angular momentum balance:

$$V = \frac{f}{2} \frac{R^2 - r^2}{r}$$

$$\frac{f}{2} \frac{dR^2}{dt} = -gr \frac{\partial \tau_\theta}{\partial P}$$

$$\rightarrow r \frac{\partial r^2}{\partial R^2} \frac{2}{f} g \frac{\partial \tau_\theta}{\partial P} \cong -\frac{\partial \psi}{\partial P},$$

$$\psi_0 = -\frac{2}{f} g \frac{\partial r^2}{\partial R^2} \tau_s = \frac{2}{f} g \frac{\partial r^2}{\partial R^2} \rho_s C_D |\mathbf{V}| V$$

Saturation entropy:

$$\frac{\partial s^*}{\partial \tau} = \frac{\Gamma_d}{\Gamma_m} \left[(M_u - M_d - w) \frac{\partial s_d}{\partial z} + \frac{\dot{Q}_{rad}}{T} \right]$$

Downdraft:

$$M_d = (1 - \varepsilon_p) M_u$$

Boundary layer entropy:

$$h_s \frac{\partial s}{\partial \tau} = C_k |\mathbf{V}| (s_0^* - s) + C_D |\mathbf{V}|^3 - (M_u - w_0)(s - s_m) + C_D r |\mathbf{V}| V \frac{\partial s}{fR\partial R}$$

Used to define M_{ueq} when > 0 ; otherwise, equation integrated for s

Relaxation equation:

$$\frac{\partial M_u}{\partial \tau} = \frac{M_{ueq} - M_u}{\tau_c}$$

Precipitation efficiency:

$$\varepsilon_p = \frac{S_m - S_{m0}}{S - S_{m0}}$$

Middle troposphere entropy:

$$\frac{\partial s_m}{\partial \tau} = \Lambda M_u (s - s_m) + \dot{Q}_{rad}$$

Radiation: $\dot{Q}_{rad} \approx -\left(s^* - s^*(t=0)\right)$

Radial diffusion added to equations

for r_b , r_t , s^* , and s_m

$$D_b = -\frac{1}{R} \frac{\partial}{\partial R} \left[r_b^2 \nu_b \frac{\partial}{\partial r_b} \left(\frac{R^2}{r_b^2} \right) \right]$$

$$D_t = -\frac{1}{R} \frac{\partial}{\partial R} \left[r_t^2 \nu_t \frac{\partial}{\partial r_t} \left(\frac{R^2}{r_t^2} \right) \right]$$

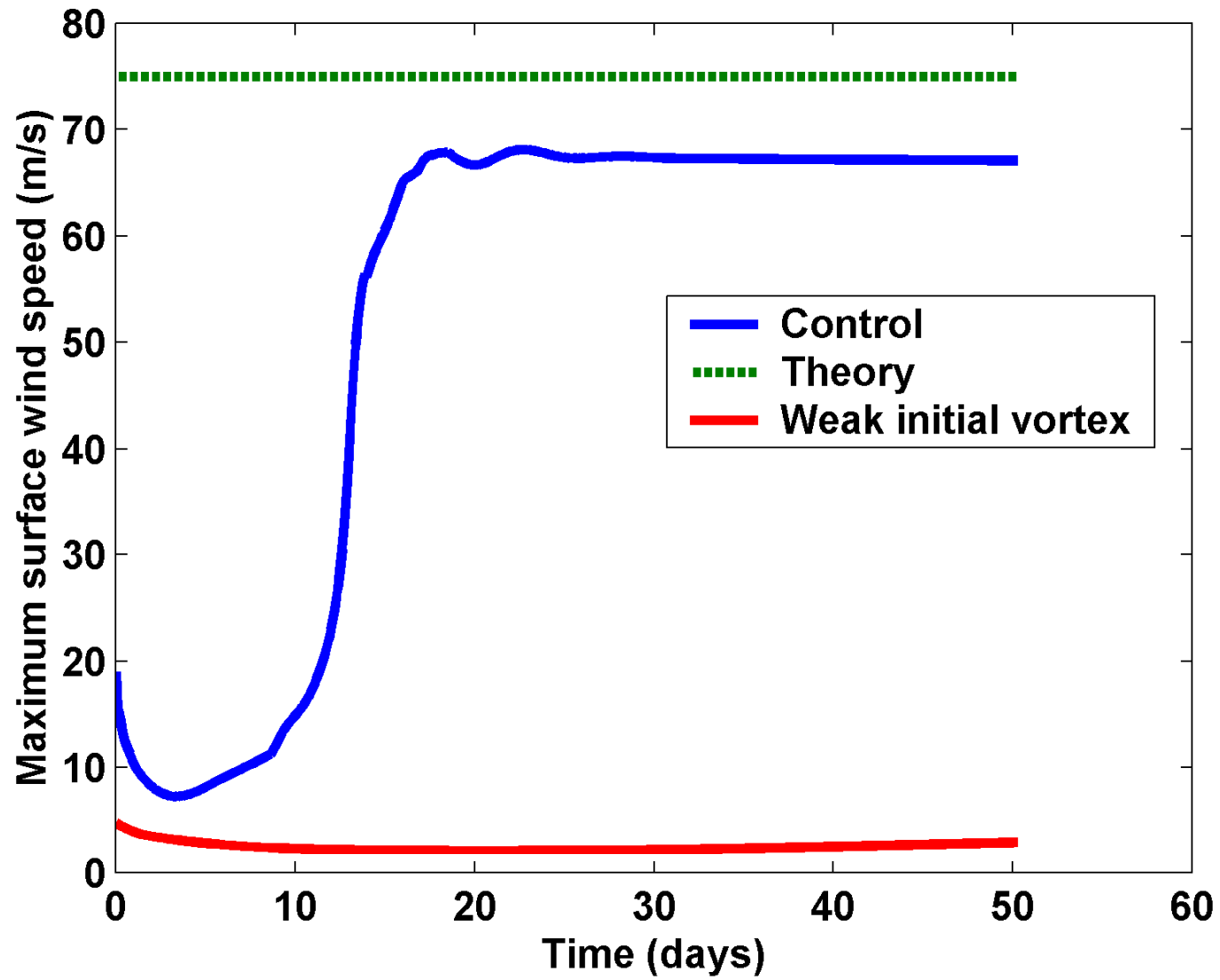
$$D_{s^*} = \frac{\partial}{\partial r_b^2} \left(r_b \nu_b \frac{\partial s^*}{\partial r_b} \right) \quad D_{s_m} = \frac{\partial}{\partial r_m^2} \left(r_m \nu_b \frac{\partial s_m}{\partial r_m} \right)$$

$$v_i = l^2 \left| r_i \frac{\partial}{\partial r_i} \left(\frac{R^2}{r_i^2} \right) \right|$$

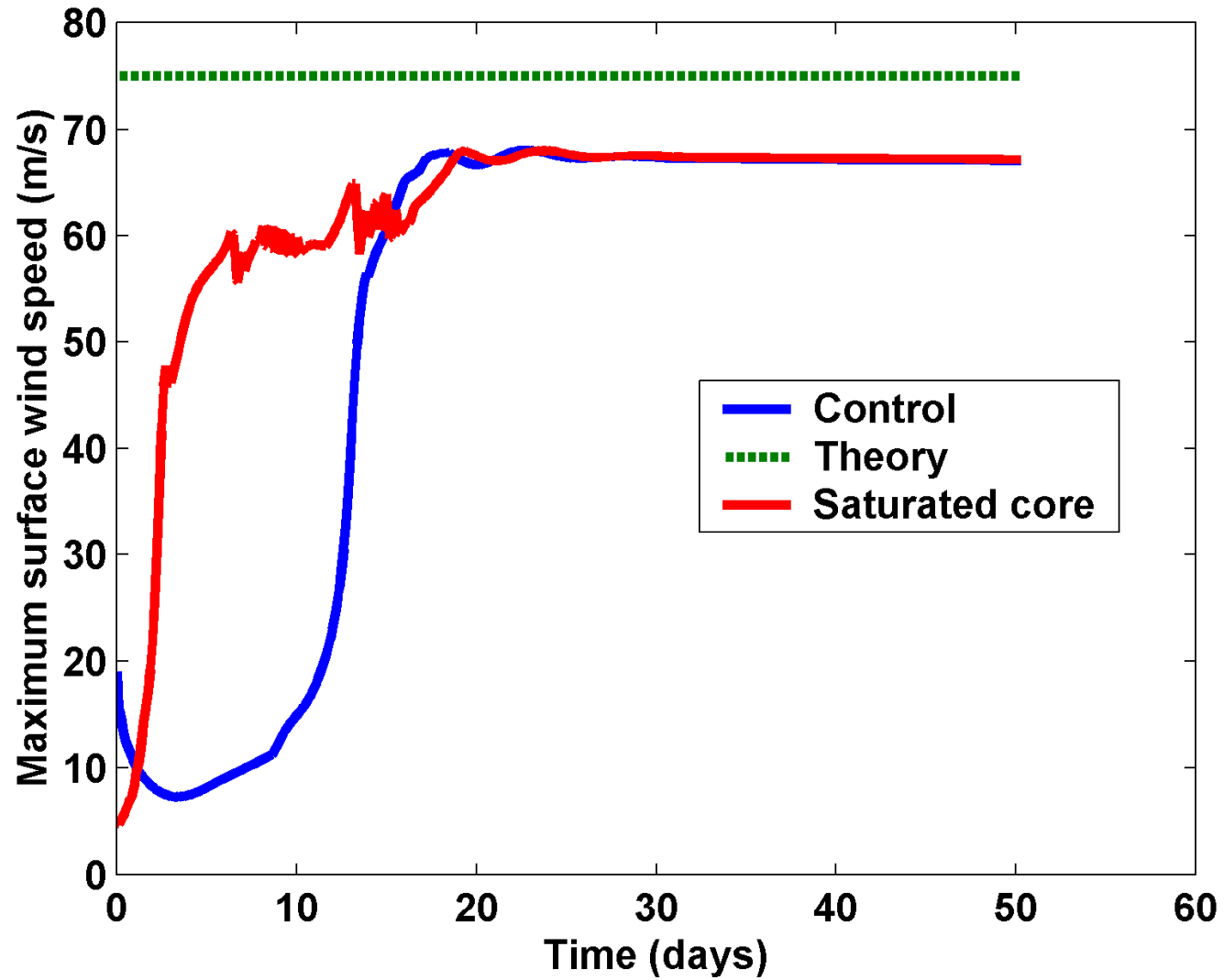
Note that surface saturation entropy depends on pressure, which is calculated from gradient wind balance using V

Complete equations summarized in Emanuel (1995), posted on course web page.

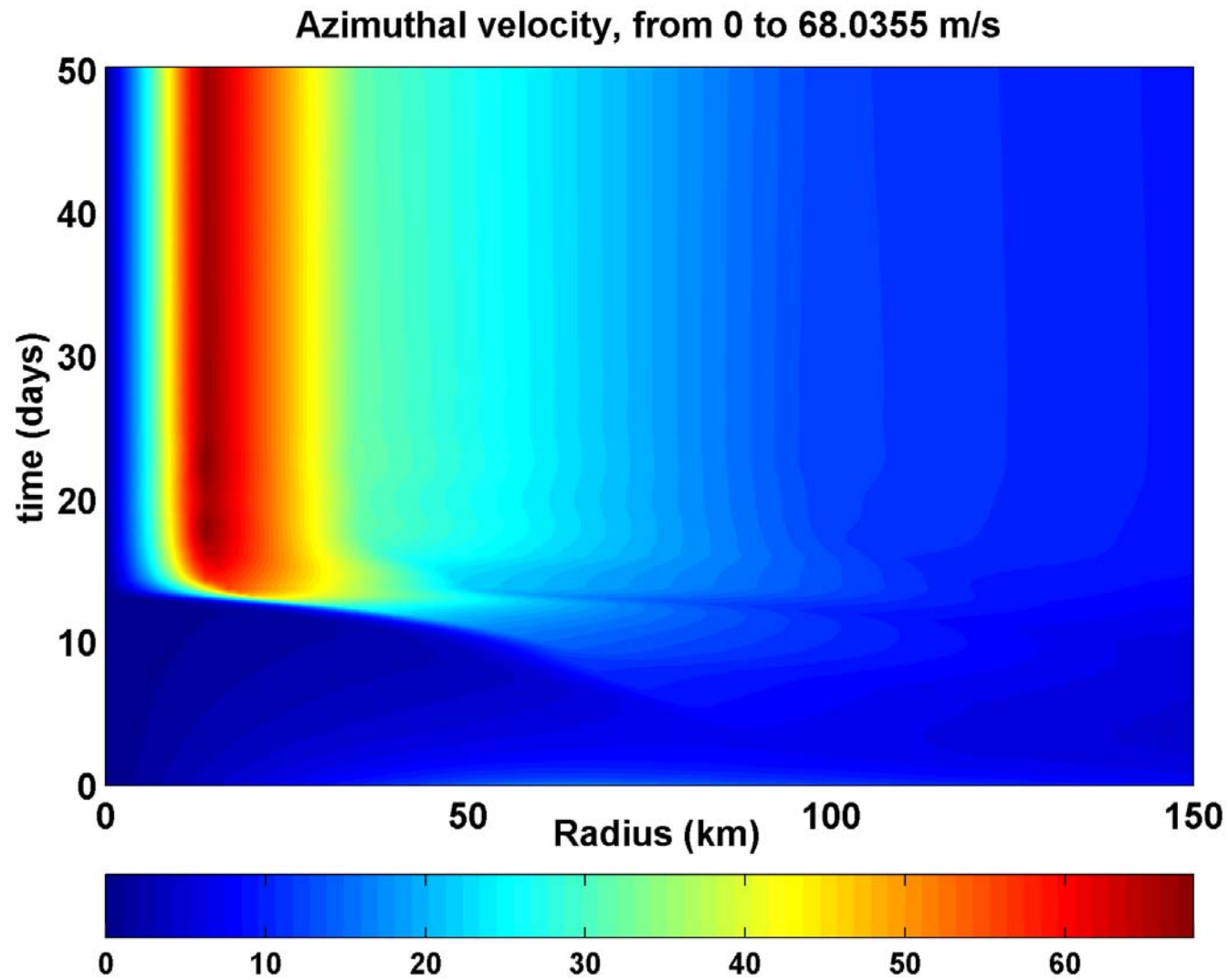
Model behavior



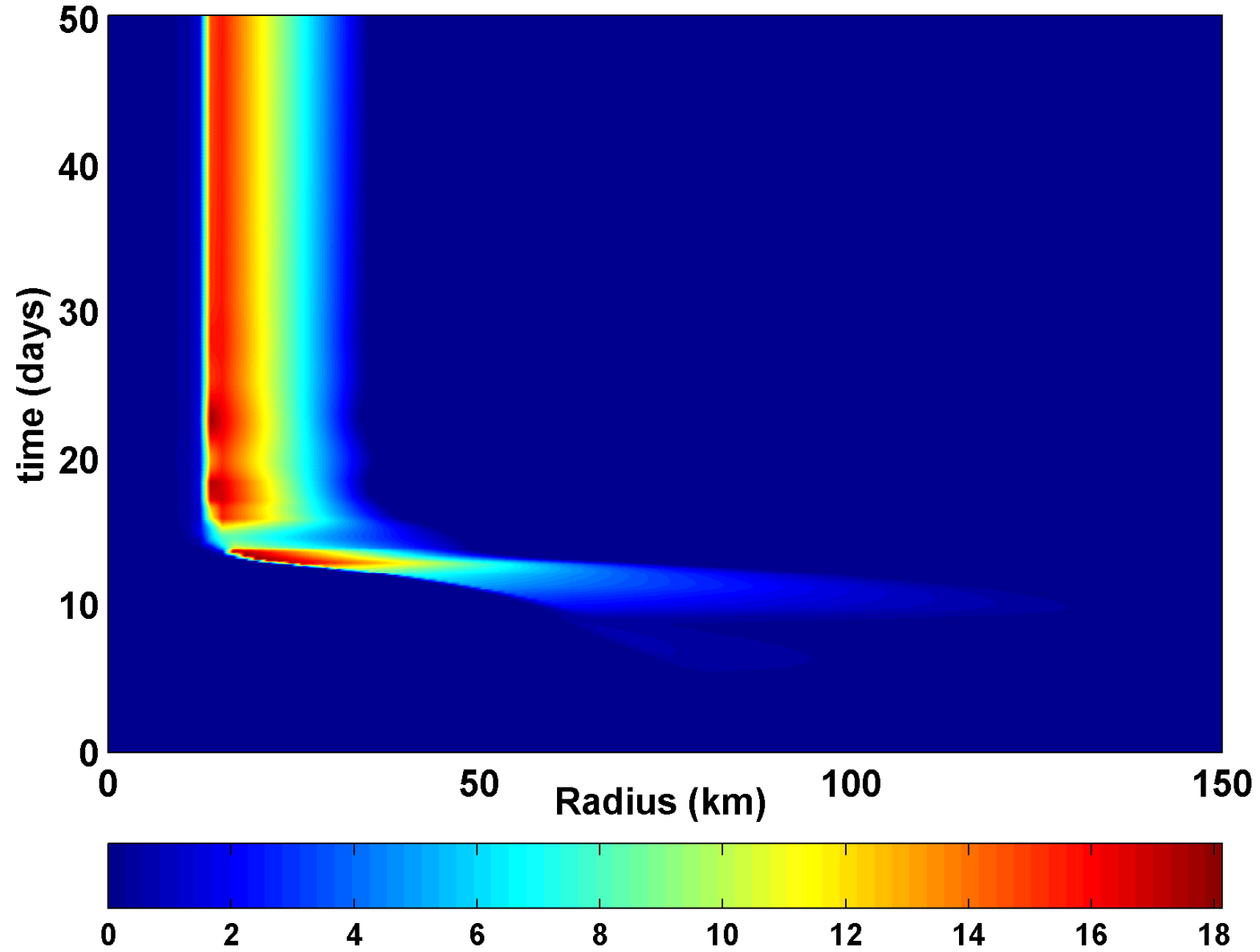
Saturate troposphere inside 100 km in initial state:



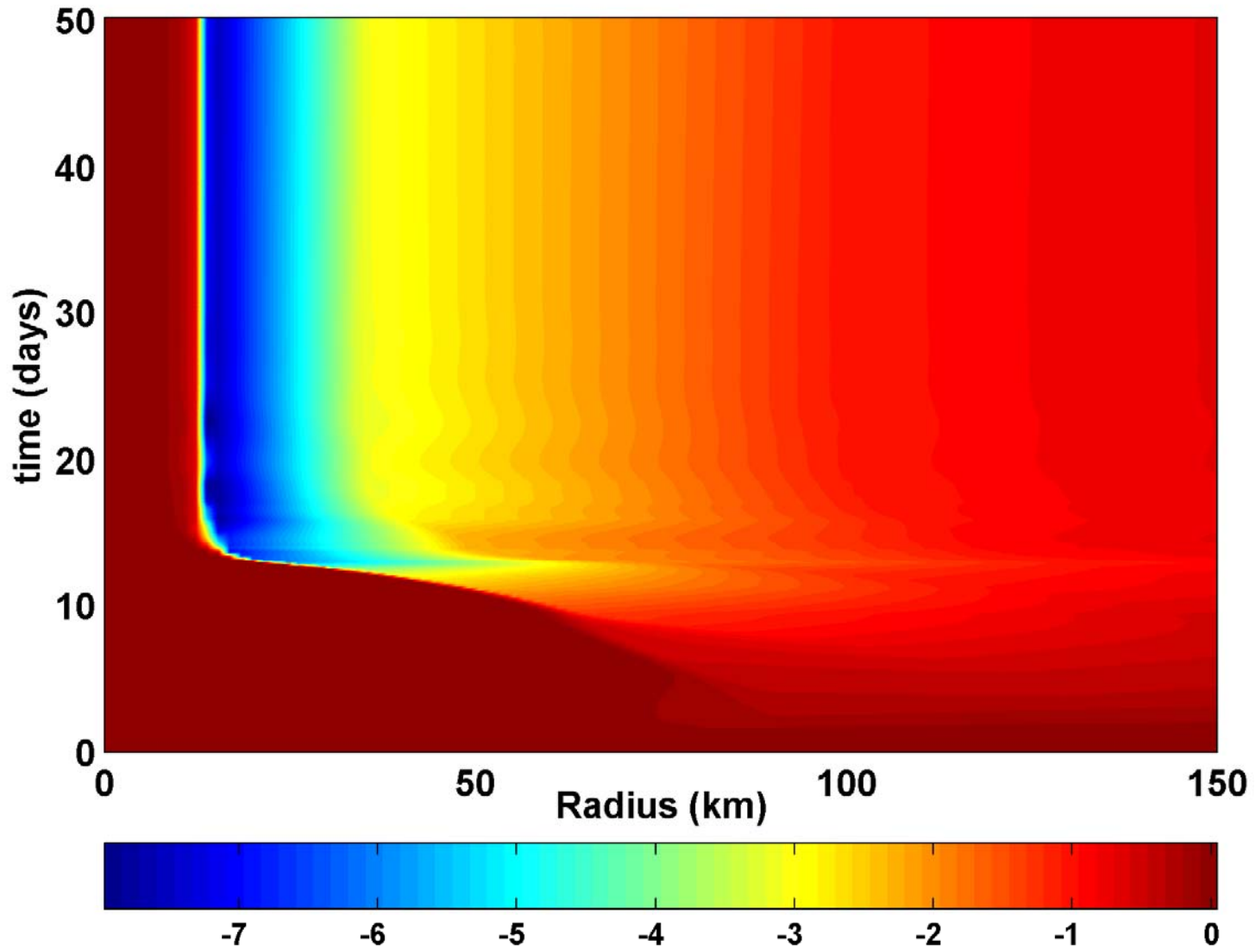
Character of control simulation

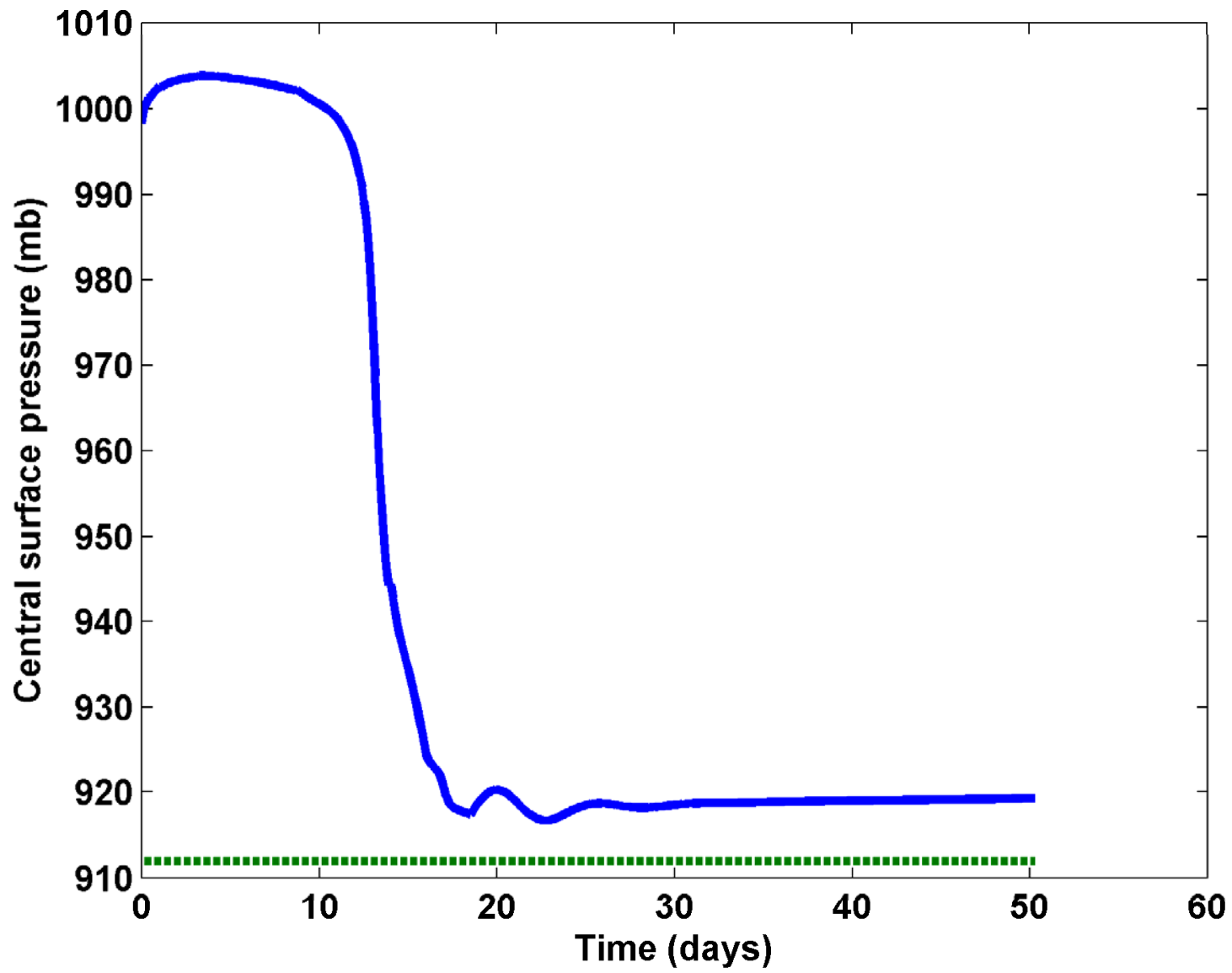


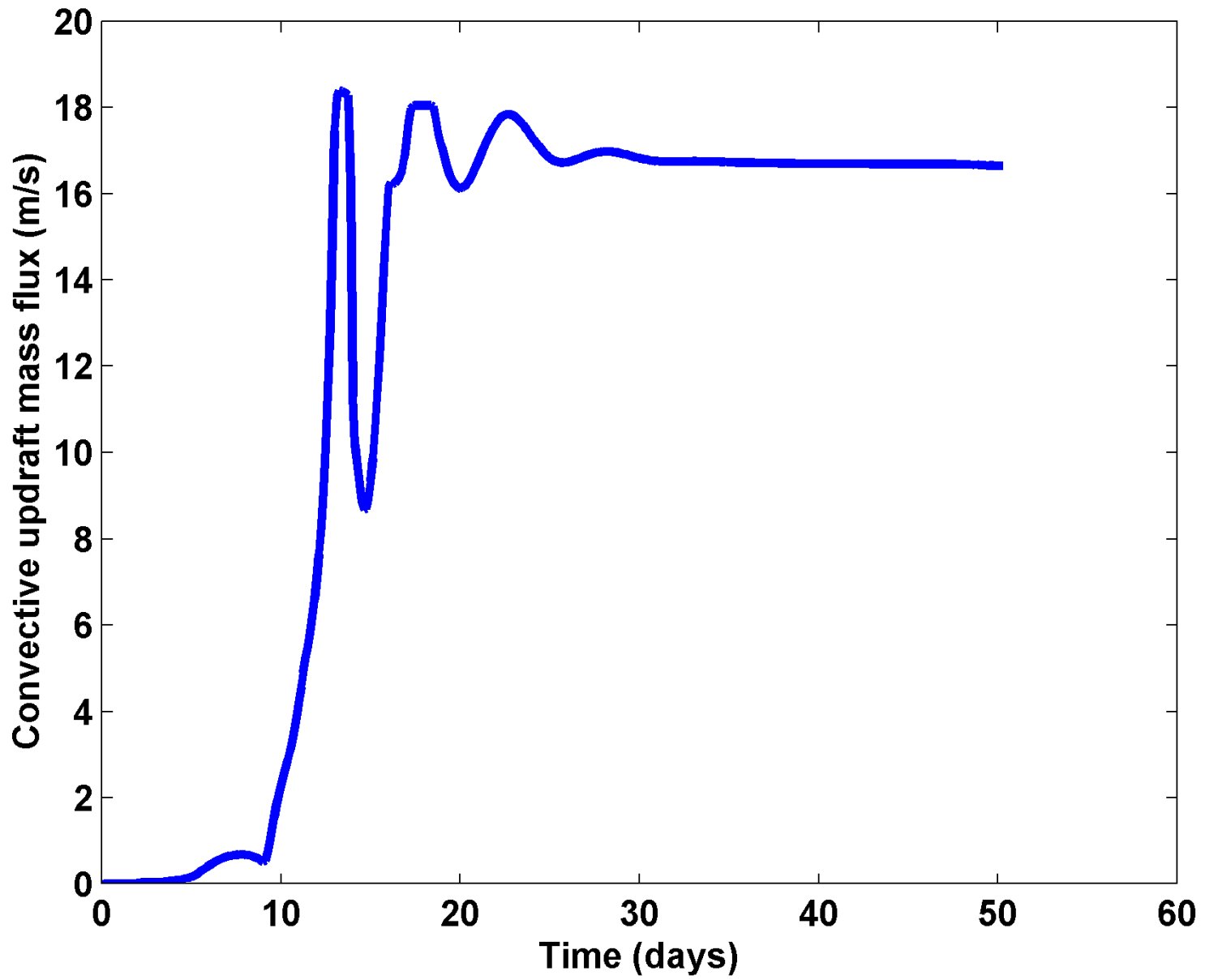
Cumulus mass flux, from 0 to 18.1277 m/s

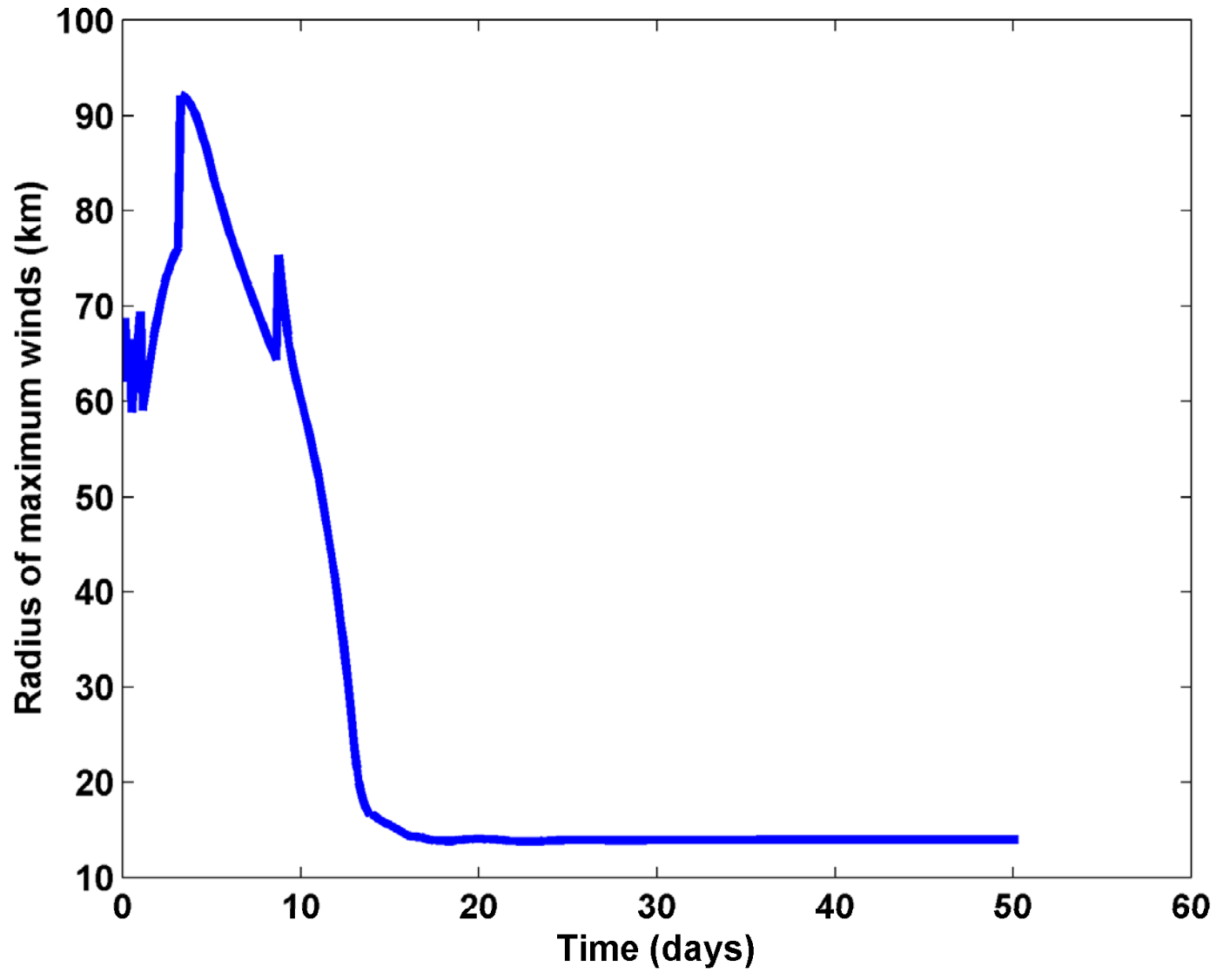


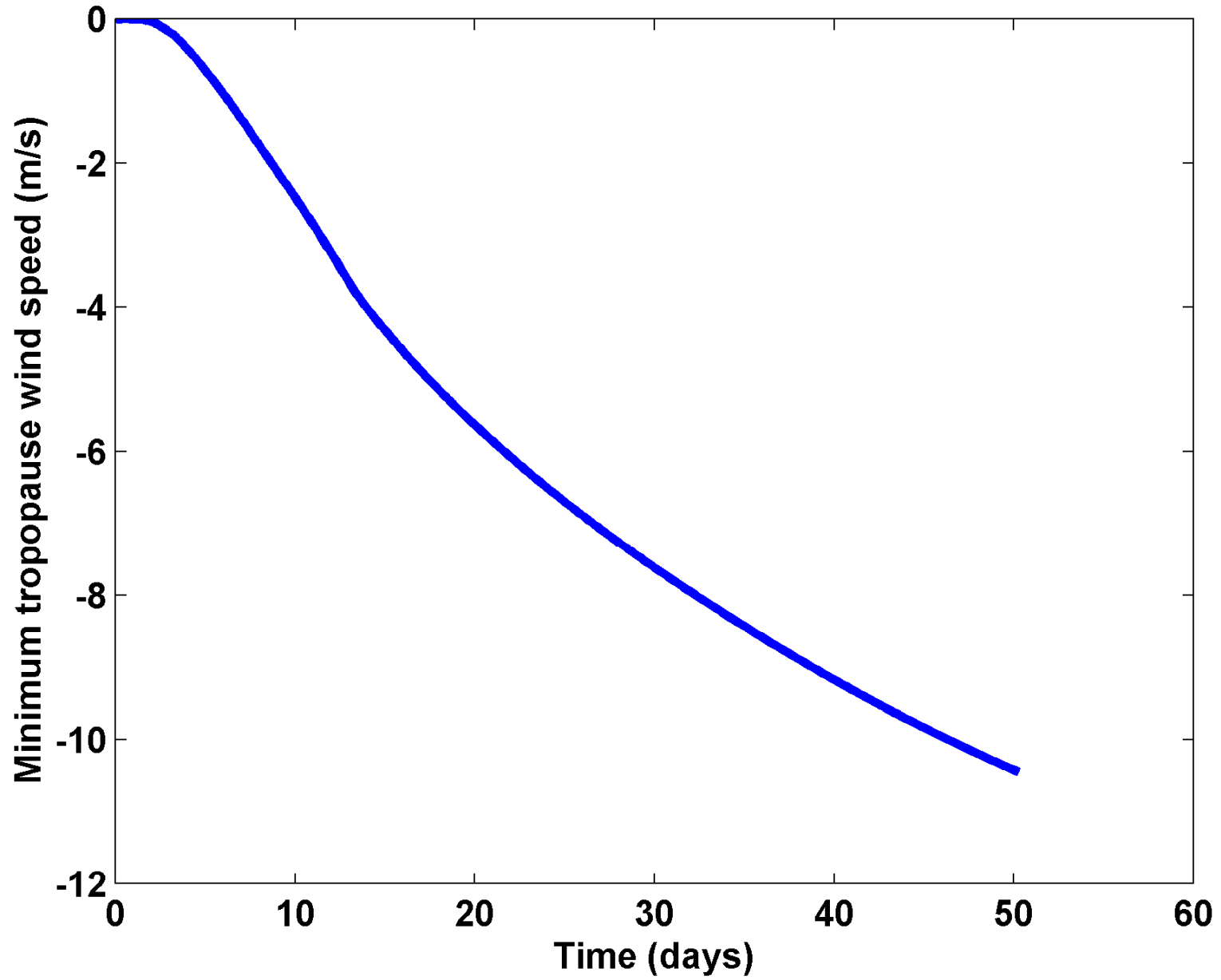
Radial velocity, from -7.9655 to 0.044 m/s

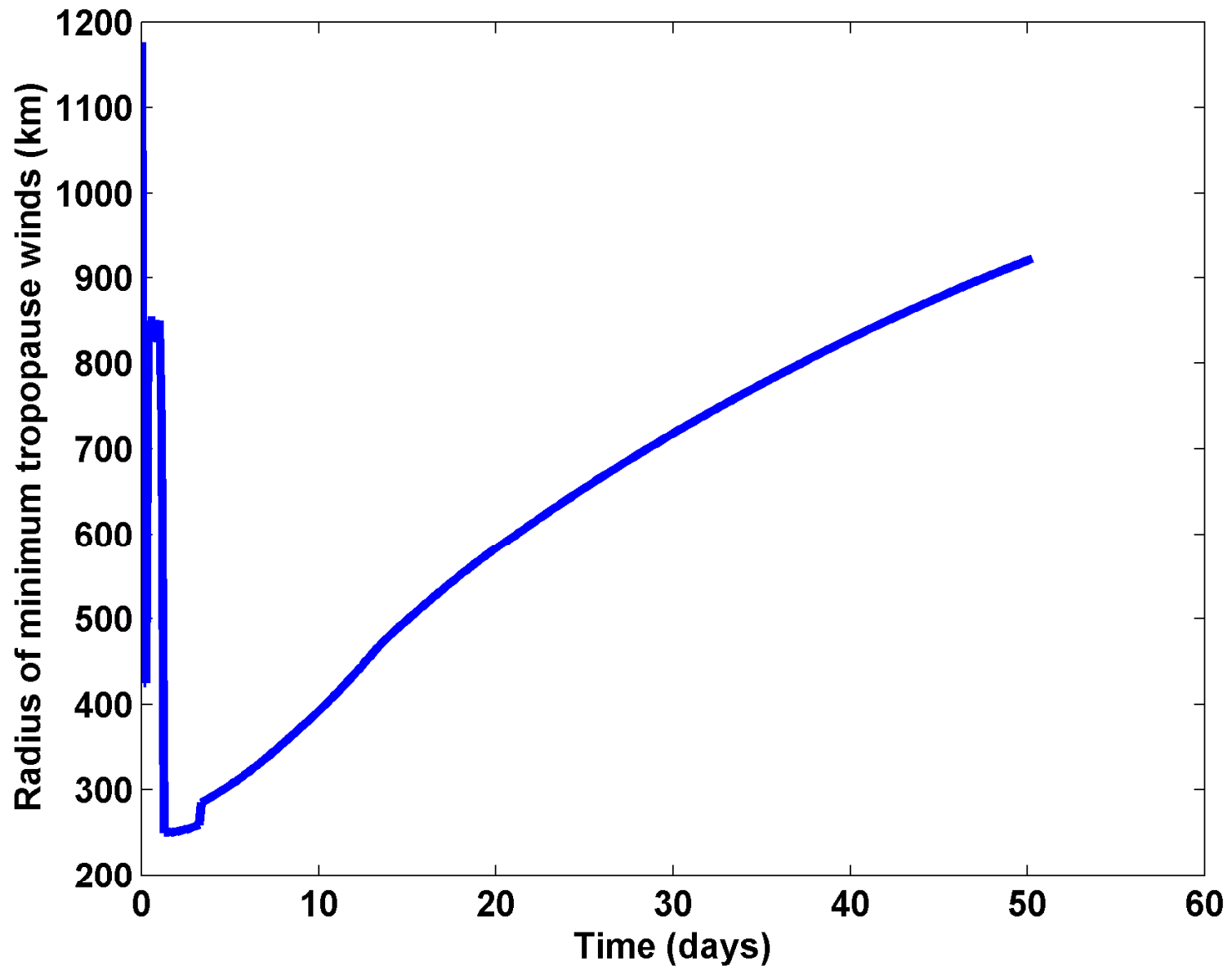




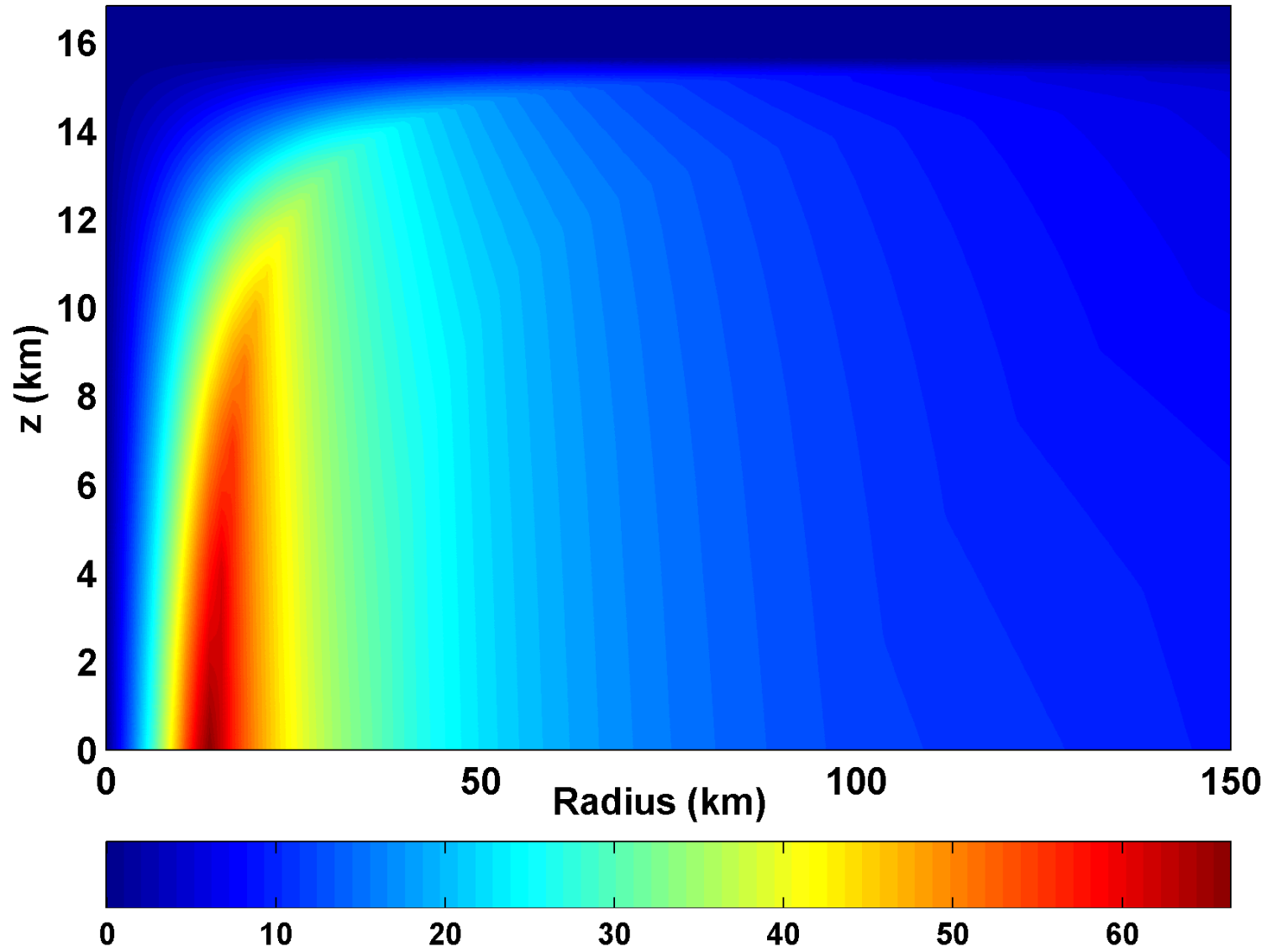




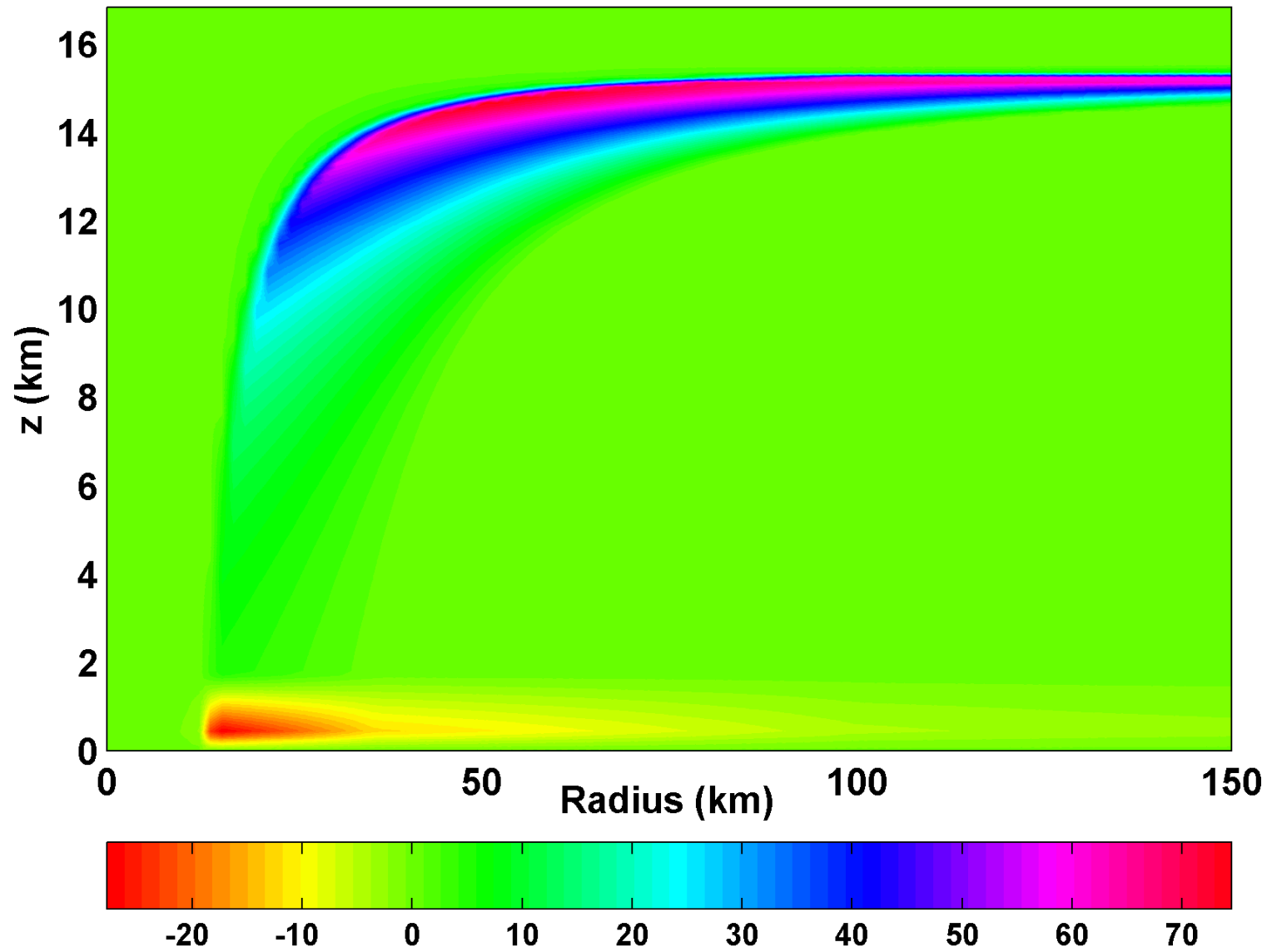




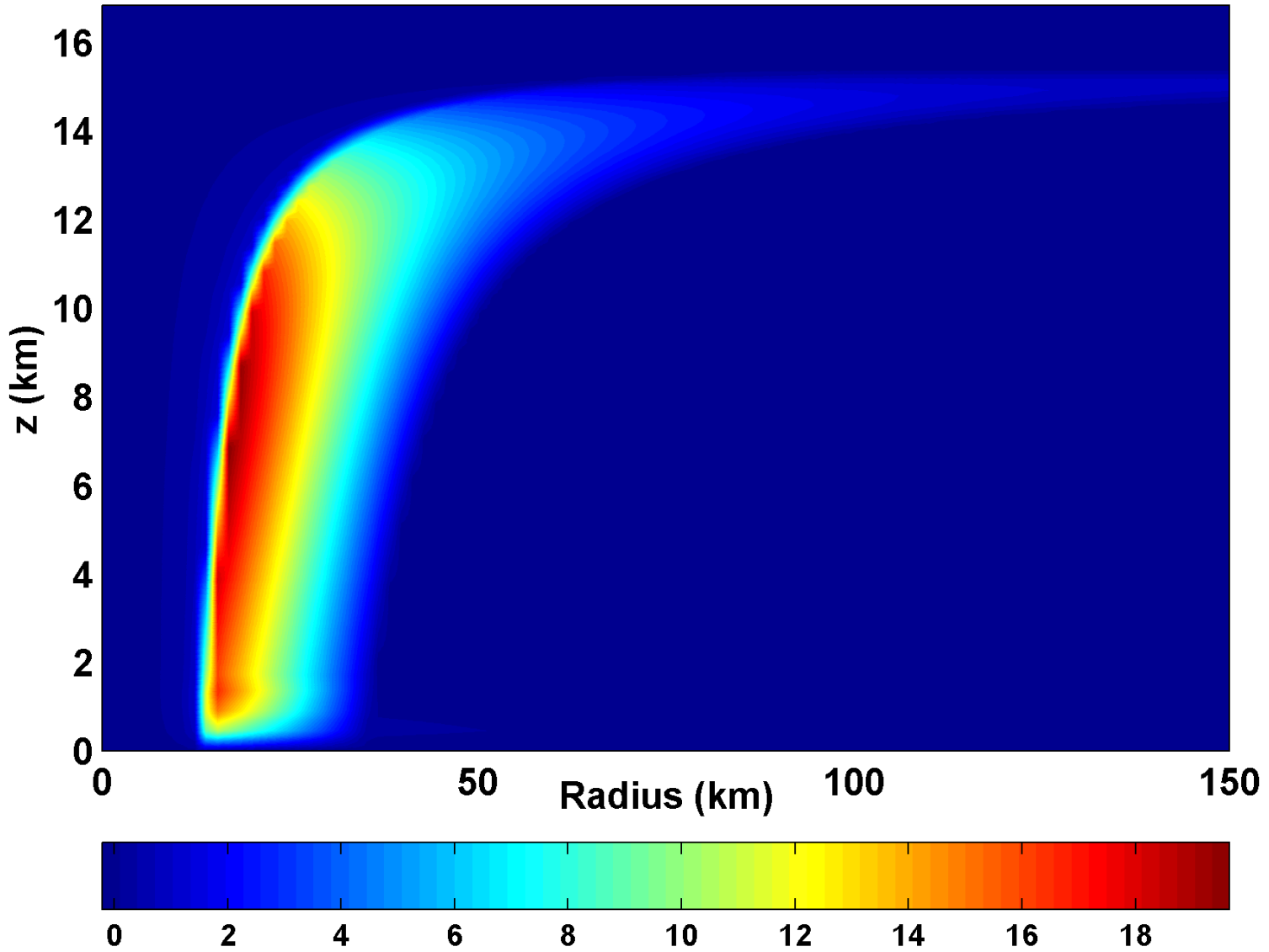
Azimuthal velocity, from -0.0423 to 66.4187 m/s



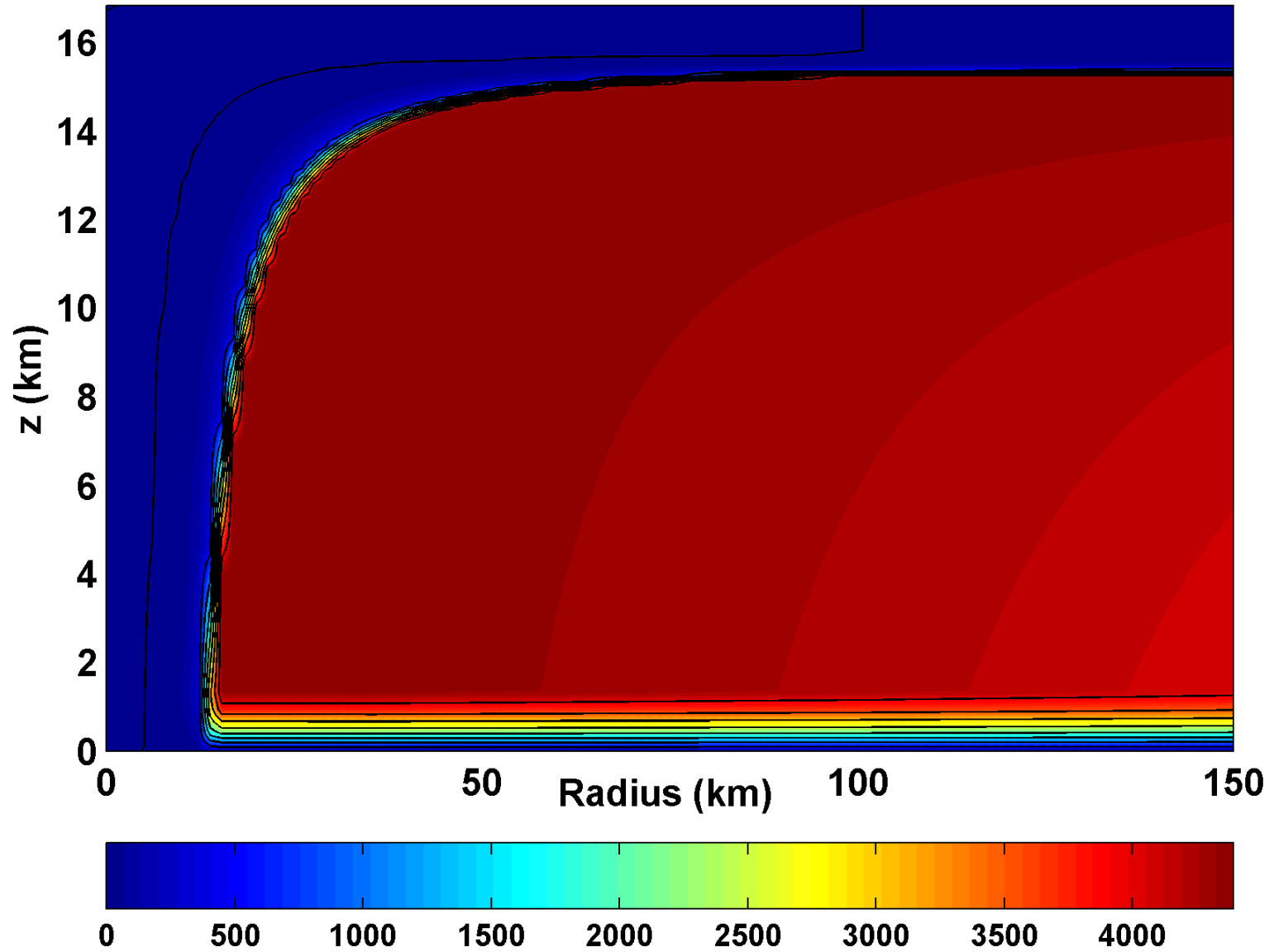
Radial velocity, from -27.7593 to 74.5129 m/s



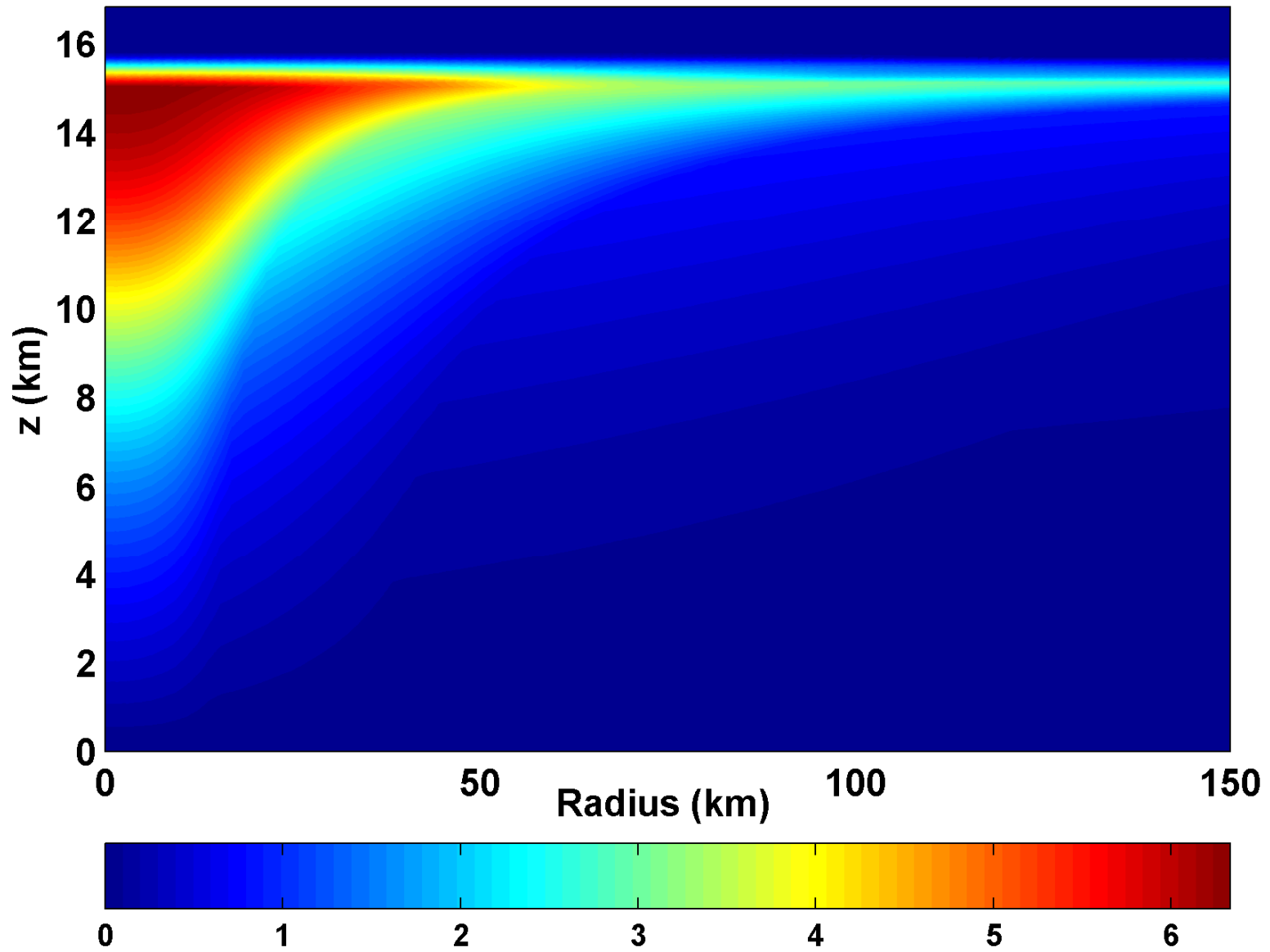
Vertical velocity, from -0.2099 to 19.6568 m/s (- values X 10)



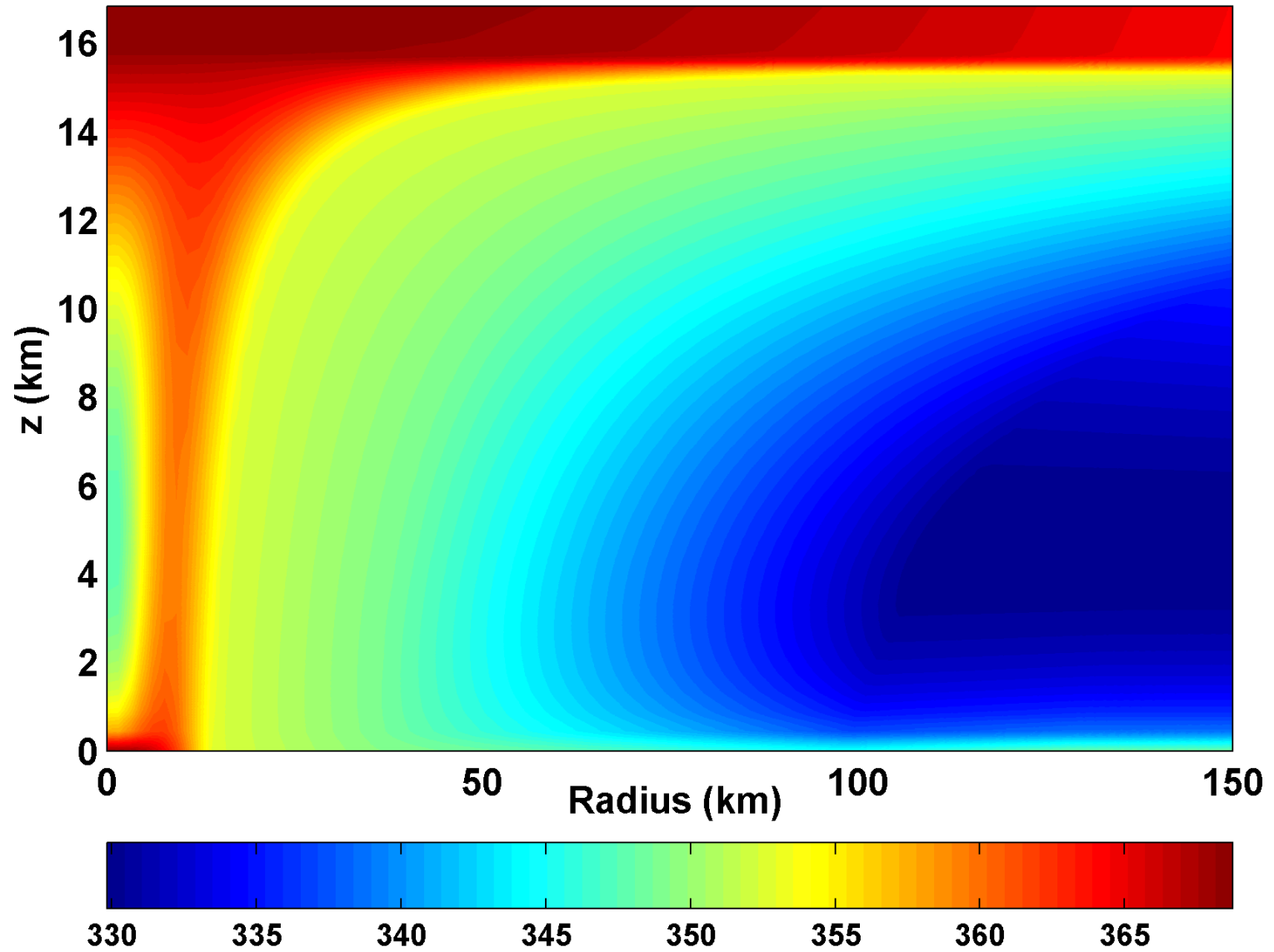
Streamfunction, from -0.8314 to 4393.2822 10^{**8} Kg/s



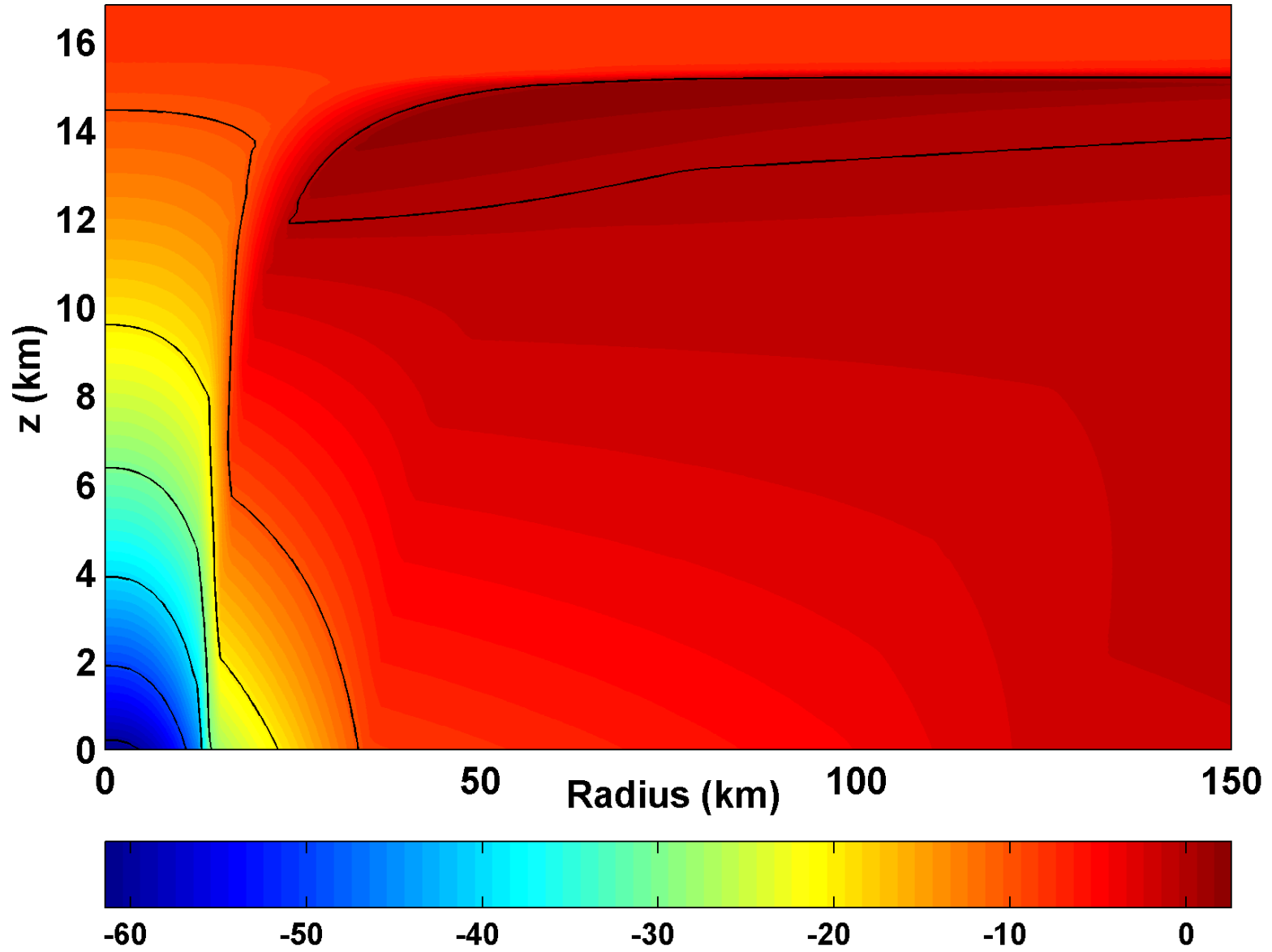
Perturbation temperature, from -0.0001 to 6.3348 K



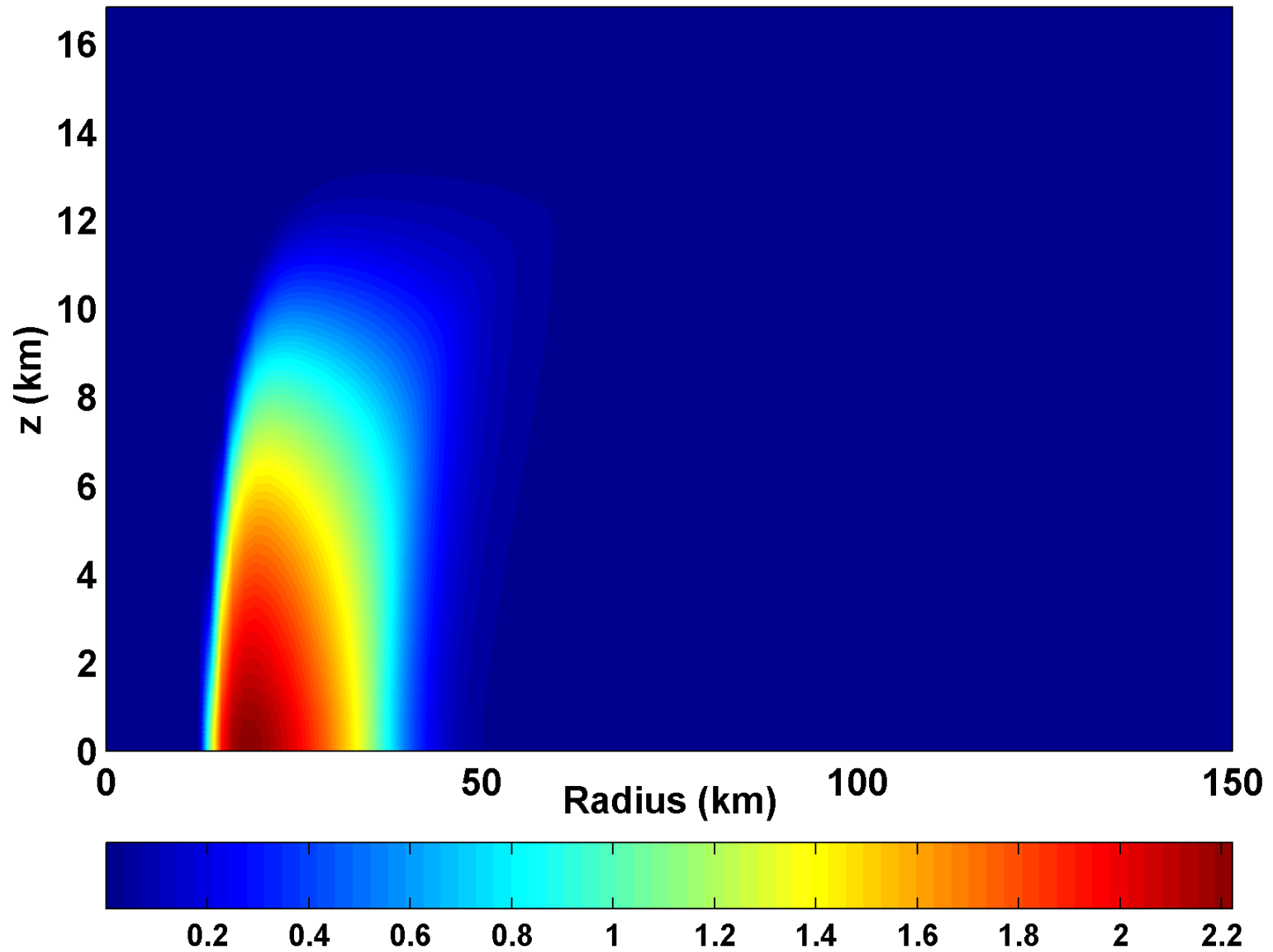
Equivalent potential temperature, from 329.8344 to 368.7422 K



Perturbation pressure, from -61.4327 to 2.5845 mb



Log of Rain water content, from 0 to 8.2261 g/Kg



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12.811 Tropical Meteorology
Spring 2011

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